

A new tidal analysis of superconducting gravity observations in Western and Central Europe

Branislav HÁBEL¹, Bruno MEURERS²

¹ Department of Theoretical Geodesy, Faculty of Civil Engineering
Slovak University of Technology in Bratislava
Radlinského 11, 813 68 Bratislava, Slovak Republic
e-mail: branislav.habel@stuba.sk

² Department of Meteorology and Geophysics
Faculty of Earth Sciences, Geography and Astronomy, University of Vienna
Althanstrasse 14, A-1090 Wien, Austria
e-mail: bruno.meurers@univie.ac.at

Abstract: Tidal analysis was applied on 1 hour gravity data acquired by a subnet of 12 superconducting gravimeters situated in western and central parts of Europe. Tidal parameters for O_1 and M_2 were adjusted. Most of the gravity time series were provided by the Global geodynamics project. Filter effects of the decimation process were carefully studied and taken into account where necessary. The ocean loading effect included in observed tidal parameters was removed using 8 ocean models with different spatial resolutions. Two different comparisons of corrected tidal parameters were performed. As a result, the accuracy of ocean loading correction and global calibration error were evaluated. For O_1 , amplitude factors are in good agreement with the hydrostatic/elastic model DDW/He while the non-hydrostatic/inelastic model DDW/NHi better describes the M_2 amplitude factors. The analysis of residual vectors allowed assessing the efficiency of the used ocean loading models which is about 81 % (O_1) and 97 % (M_2).

Key words: tidal parameters, ocean tides, body tide models, Global geodynamics project

1. Introduction

Relative and absolute gravity measurements are inherently related to the modeling of tides at a certain point on the Earth's surface. Calculation of tidal effects is based on either theoretical tidal models or results of the tidal analysis applied to gravity time series. The first approach is depending on

the accuracy of used theoretical models and it is questionable how these models describe the tidal deformation of the Earth or the indirect effect of the ocean tides. On the other hand, the determination of observed tidal parameters allows computing a total tidal signal at a given site including all deformations and potential changes due to the tides. However, this approach suffers from inaccurate scale factors of relative gravimeters.

Various authors have been dealing with the determination and comparison of tidal parameters derived from relative gravity measurements of both classical spring and superconducting gravimeters (SG), e.g. *Baker et al. (1991)*; *Baker et al. (1996)*; *Ducarme et al. (2002)*; *Baker and Bos (2003)*; *Ducarme et al. (2009)*. This paper is focused on SG gravity observations based on a dense network of stations in western and central Europe with new and long time series of at least 3 years. SGs are characterized by a small and regular instrumental drift and therefore enable continuous observation of very small changes in gravity acceleration with high accuracy and long-time stability. This makes them an ideal instrument for high-precise observation of tidal signals within a wide frequency band, including the long-period component. However, the knowledge of a precise transfer function of the gravimeter is mandatory and can be the main limitation in increasing the accuracy of tidal parameters. The determination of the transfer function is commonly carried out by different methods. Regarding the amplitude part, comparison with absolute gravity observations is the most widely used method at present. The calibration factor can be determined at the 0.1 % accuracy level using FG5 absolute gravimeters (*Francis, 1997*). The instrumental time lag of the data acquisition system has to be defined independently. Precision of better than 0.01 s can be achieved (*Van Camp et al., 2000*). From the perspective of tidal analysis, preprocessing of gravity observations plays another important role, e.g. data resampling and treatment of instrumental artifacts, earthquakes or other disturbances. Such artifacts as well as calibration errors can corrupt the tidal analysis results. They are not investigated in this paper. Certainly, their presence in tidal signal degrades the quality of analysis.

SGs are the basic instrumentation of the Global geodynamics project (GGP) since 1997 (*Crossley et al., 1999*). GGP is a worldwide network monitoring gravity changes with high accuracy for the purpose of geodynamics research. Nowadays thanks to new SGs integrated into GGP, more

precise calibration, new ocean loading models and longer gravity time series, it is possible to perform new comparisons of tidal parameters. It enables to evaluate the calibration error and the accuracy of ocean loading correction or to validate body tide models.

In this paper, we assess the efficiency of ocean loading corrections with respect to different ocean loading models and the hydrostatic and non-hydrostatic body tide models by *Dehant et al. (1999)*. Also, we display the spatial dependence of the corrected tidal parameters on the location. The amplitude factors derived from the body tide models shows only weak latitude dependence. Within the investigated area, they increase only by 0.04 ‰ (O_1) and 0.07 ‰ (M_2) respectively from South to North. This helps to identify stations with systematic calibration errors on one hand, and station dependent deficiencies of the ocean loading correction on the other.

2. Tidal parameters theory

Tidal analysis consists in determination of observed amplitude factors δ and phase differences α for a given group of tidal waves combining constituents with similar frequencies which cannot be separated due to the insufficient length of the time series (e.g. *Melchior et al., 1996*). The dimensionless amplitude factor δ is defined as the ratio A/A_{th} where A is the observed amplitude of a given tidal wave derived from tidal analysis and A_{th} is the theoretical amplitude related to the astronomical tides for an ocean-less rigid Earth (e.g. *Melchior, 1978*). The ratio A/A_{th} reflects additional gravity changes due to tidal deformation of the Earth's body and the ocean loading effect. Theoretical tidal parameters are provided by calculating the tidal response of different Earth's models (*Dehant et al., 1999; Mathews, 2001*). The theoretical amplitude factors δ_{th} express only the tidal deformation of the Earth without oceans. Theoretical phase differences α_{th} are negligible with respect to their magnitude. *Dehant and Zschau (1989)* conclude that an effect of mantle inelasticity is of the order of -0.005° for frequencies between the semi-diurnal and the monthly tides.

Observed tidal parameters are components of the observed tidal vector $\mathbf{A}(A, \alpha) = \mathbf{A}(\delta \cdot A_{th}, \alpha)$ which is plotted in a polar coordinate system in Fig. 1. The theoretical tidal vector $\mathbf{R}(\delta_{th} \cdot A_{th}, 0)$ is defined in a similar way.

Figure 1 also shows the ocean loading vector \mathbf{L} (L, λ) with amplitude L and phase λ related to the local meridian. The ocean loading effect depends on the distance between gravimeter and sea or ocean, and is stronger near the coastlines. Comparison of tidal parameters of different stations can only be done after removing the ocean loading effect resulting to the corrected tidal vector \mathbf{A}_c (A_c, α_c) with corrected amplitude A_c and phase difference α_c (Ducarme *et al.*, 2009) according to Eq. (1):

$$\mathbf{A}_c(\delta_c \cdot A_{th}, \alpha_c) = \mathbf{A} - \mathbf{L}, \quad (1)$$

where δ_c is corrected amplitude factor.

For completeness, the residual vectors \mathbf{B} and \mathbf{X} are also shown in Fig. 1. They are derived from comparison of observed and theoretical tidal vectors:

$$\mathbf{B}(B, \beta) = \mathbf{A} - \mathbf{R}, \quad (2)$$

$$\mathbf{X}(X, \chi) = \mathbf{B} - \mathbf{L}. \quad (3)$$

If the ocean loading and body tide models are perfect and describe the real mass redistribution correctly, then the components of the residual vector \mathbf{X} reflect the errors of the instrumental transfer function or instrumental noise (e.g. Melchior, 1994). On the other hand, if the transfer function is

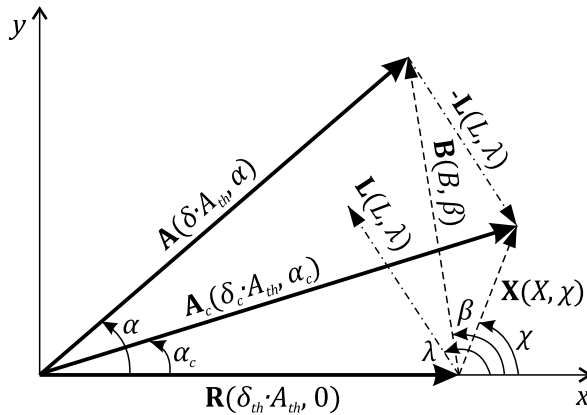


Fig. 1. Illustration scheme of observed tidal vector $\mathbf{A}(A, \alpha)$, theoretical tidal vector $\mathbf{R}(\delta_{th} \cdot A_{th}, 0)$, ocean loading vector $\mathbf{L}(L, \lambda)$, observed residual vector $\mathbf{B}(B, \beta)$ and remaining residual vector $\mathbf{X}(X, \chi)$ in polar coordinate system (Neumeier *et al.*, 2005).

determined correctly and the body tide model is perfect, then \mathbf{X} reflects deficiencies of the ocean loading model. Actually, these assumptions are never perfectly true, i.e. calibration and ocean loading model errors as well as the choice of a theoretical body tide model contribute to both components of the residual vectors. Nevertheless, the analysis of residual vectors allows studying deficiencies of ocean loading models, Earth's models and SG calibration.

3. Data processing and tidal parameters estimation

A subnet of 12 SGs is situated in western and central parts of Europe. The location of the stations used in this study is shown in Fig. 2. Distances from site to site vary between 50 km and 1000 km. Consequently, the tidal parameters are affected by ocean tides in different ways. Station details are listed in Table 1. Earlier type SGs marked by “T” (tidal) were recording in Brussels and Potsdam. The other stations have been equipped with newer generation SGs like “CT” (compact tidal), “CD” (double sphere) and “OSG” (observatory SG). In case of dual sphere SGs only the gravity data from the lower sphere (L) was used. The determination of calibration factors has been mostly done by co-located gravity observations using FG5 absolute gravimeters. Older types SG T-018 and T-003 were calibrated by using relative spring gravimeter Lacoste&Romberg D02 in Potsdam and Scintrex CG3M in Brussels (*Ducarme et al., 2009*).

3.1 Data processing

Data are provided by the Information System and Data Center (ISDC) of the German Research Centre for Geosciences – GFZ¹ in the frame of GGP. In our study, we used the 1 hour gravity data processed by the International Center for Earth Tides (ICET) except for the stations Conrad (CO), Vienna (VI) and Walferdange (WA) that were obtained directly from their operators (B. Meurers, O. Francis). Other operators have been contacted to check the calibration parameters (Membach, Moxa, Pecny and Potsdam). For some stations, incompatibility with the GGP database was found, and

¹ <http://isdc.gfz-potsdam.de/>

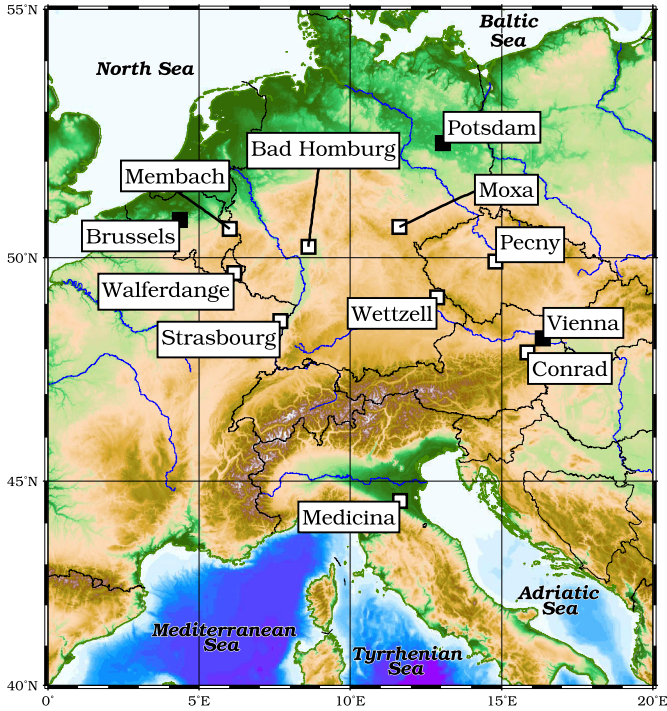


Fig. 2. Network of SG stations used in this study located in western and central Europe. Sites in white are currently in operation.

updated calibration parameters were considered. The original 1 s gravity data of VI and CO was first re-sampled to 1 min data by applying the Chebyshev filter G1S1M and then decimated to 1 hour samples by applying the filter G1M1H². Both filters are provided by GGP and recommended due to its negligible amplitude attenuation within all tidal bands. These filters are implemented in ETERNA 3.40 software package (Wenzel, 1996). ICET decimates the 1 min samples of ISDC by applying the TSOFT software (Van Camp and Vauterin, 2005). Before resampling to 1 hour data, a LSQ low-pass filter with length of 16 hours and cut-off frequency of 12 cpd (cycles per day) is used (B. Ducarme, pers. comm.). The 1 min data of WA was re-sampled accordingly (O. Francis, pers. comm.). However, fil-

² <http://www.eas.slu.edu/GGP/ggpfilters.html>

Table 1. Selected SG stations located in western and central Europe

SG	Location	Latitude (°N)	Longitude (°E)	Height (m)	SG type	Length of records (year)
BE	Brussels (<i>Belgium</i>)	50.7986	4.3581	100.0	T-003	12.28
BH	Bad Homburg (<i>Germany</i>)	50.2285	8.6113	190.0	OSG-044	3.82
CO	Conrad (<i>Austria</i>)	47.9283	15.8598	1045.0	CT-025	3.72
MB	Membach (<i>Belgium</i>)	50.6093	6.0066	250.0	CT-021	9.72
MC	Medicina (<i>Italy</i>)	44.5219	11.6450	28.0	CT-023	3.50
MO	Moxa (<i>Germany</i>)	50.6447	11.6156	455.0	CD-034L	7.73
PE	Pecny (<i>Czech</i>)	49.9138	14.7856	534.6	OSG-050	3.07
PO	Potsdam (<i>Germany</i>)	52.3806	13.0682	81.0	T-018	6.11
ST	Strasbourg (<i>France</i>)	48.6217	7.6838	180.0	CT-026	7.79
VI	Vienna (<i>Austria</i>)	48.2489	16.3565	192.0	CT-025	11.74
WA	Walferdange (<i>Luxembury</i>)	49.6647	6.1528	295.0	OSG-040	5.02
WE	Wetzell (<i>Germany</i>)	49.1440	12.8780	613.7	CD-029L	3.46

tering can affect the amplitudes and phases of tidal constituents depending on their frequencies and can systematically influence the tidal parameters. This should be carefully studied and corrected, if required.

The frequency response of a filter can be investigated by its transfer function. For that purpose, the filter is applied to a Heaviside step function with 1 min sampling. Because of their symmetrical properties, we do not expect any phase distortion. The amplitude responses of the LSQ and G1M1H low-pass filters are given in Table 2 for the principal tidal constituents. The LSQ filter shows a substantial amplitude attenuation of -0.50% in the semi-diurnal band while the phase distortion is almost zero as expected. All observed amplitude factors were, therefore, corrected accordingly.

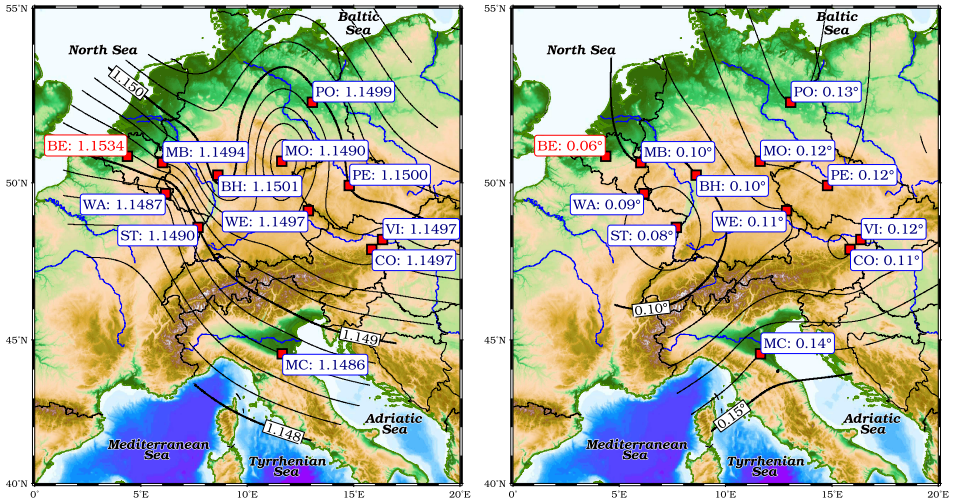
3.2 Tidal parameters estimation

The estimation of observed tidal parameters, amplitude factors δ and phase differences α , was carried out applying the ETERNA 3.40 software package (Wenzel, 1996) on 1 hour gravity and atmospheric pressure observations. Filter effects of the decimation process was taken into account where necessary (see section 3.1). The adjusted tidal parameters for O_1 and M_2 are

Table 2. Attenuation of principal tidal amplitudes in ‰ due to LSQ low-pass (16 h, 12 cpd) and G1M1H filter

Tidal wave	Frequency (cpd)	Attenuation		Tidal wave	Frequency (cpd)	Attenuation	
		LSQ	G1M1H			LSQ	G1M1H
Q1	0.893244	-0.03	0.00	M2	1.932274	-0.49	0.00
O1	0.929536	-0.04	0.00	S2	2.000000	-0.51	0.00
P1	0.997262	-0.06	0.00	K2	2.005476	-0.51	0.00
K1	1.002738	-0.06	0.00	M3	2.898410	-0.06	0.00
N2	1.895982	-0.48	0.00	M4	3.864547	0.57	-0.14

shown in Figs. 3 and 4. The theoretical amplitude of both constituents is of same order of magnitude ($\sim 300 \text{ nm.s}^{-2}$) in mid-latitudes but they are differently affected by the ocean loading effect. In both cases, isolines of tidal parameters reveal the strong dependence on the ocean loading effect as expected and slightly on the latitude. A distinct anomalous feature is clearly visible around the stations Moxa and Bad Homburg suggesting a calibration problem. A strong anomaly is obtained for Brussels too, probably due to its vicinity to the coastline, but certainly also due to imperfect calibration and other technical issues. Calibration at Brussels has never been


 Fig. 3. Observed amplitude factors (left) and phase differences (right) for O_1 .

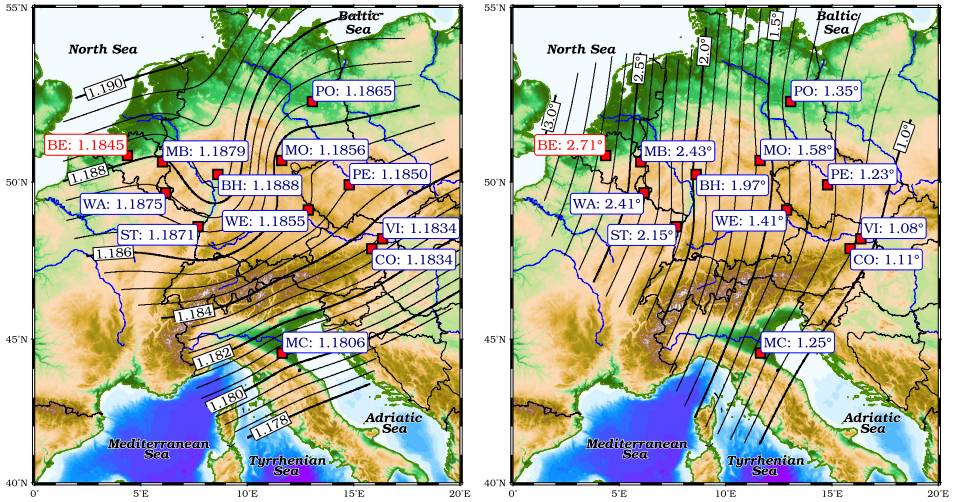


Fig. 4. Observed amplitude factors (left) and phase differences (right) for M_2 .

done by comparing with absolute gravimeters, the phase lag has never been determined. *Ducarme and Somerhausen (1993)* derived from co-located observations with a Scintrex relative gravimeter, that the SG scale factor is too high by 2%. Therefore, this station was not used for the isoline construction. Detailed analysis can be done after correcting for ocean loading effects.

4. Ocean tide models

As mentioned earlier, the gravity observations include an ocean loading effect which has to be removed before comparison of observed tidal parameters. It is caused by redistribution of ocean masses and deformation due to the load. The effect on gravity can be modeled using global ocean tide models and the Green's function approach (e.g. *Neumeyer, 2010*). Ocean loading corrections have been based on load vectors provided by M.S. Bos and G.H. Scherneck website³. This service enables selecting different ocean

³ <http://holt.oso.chalmers.se/loading/>

models and considering different types of correction.

To eliminate the influence of ocean tides on observed amplitude factors δ and phase differences α , we have used eight models with a different resolution: $0.5 \times 0.5^\circ$ – CSR4.0 (Eanes, 1994), GOT4.7 (Ray, 1999); $0.25 \times 0.25^\circ$ – TPXO7.2 (Egbert and Erofeeva, 2002); $0.125 \times 0.125^\circ$ – FES2004 (Lyard et al., 2006), EOT08a (Savcenko and Bosch, 2008), DTU10 (Cheng and Andersen, 2010), HAMTIDE (Taguchi et al., 2010) and EOT11a (Savcenko and Bosch, 2011). Each of these models provides the amplitude L and phase λ_G (with regard to Greenwich) of ocean loading vector \mathbf{L} for 11 principal tidal waves. These constituents are interpolated from the respective model at the point of computation. Then, the ocean loading phase λ related to the local meridian can be calculated according to (Francis and Melchior, 1996)

$$\lambda = -(m \cdot LONG + \lambda_G), \quad (4)$$

where $LONG$ is a longitude and m expresses the tidal band (0 for long-, 1 for diurnal- and 2 for semi-diurnal band).

Averaged amplitudes and phases of ocean loading vectors corresponding to eight above-mentioned models are shown in Figs. 5 and 6. The SG stations are sorted in alphabetical order. We focused on the O_1 and M_2 constituents only. Standard deviations reflect discrepancies between used models at the SG sites. In the study area, the O_1 load amplitude varies

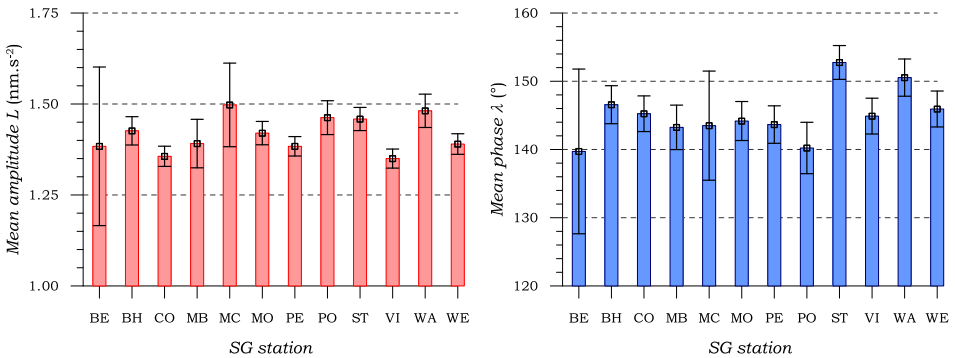


Fig. 5. Mean amplitudes (left) and phases (right) of O_1 ocean loading vector from the model predictions at different SG stations. Error bars show discrepancies between used ocean loading models.

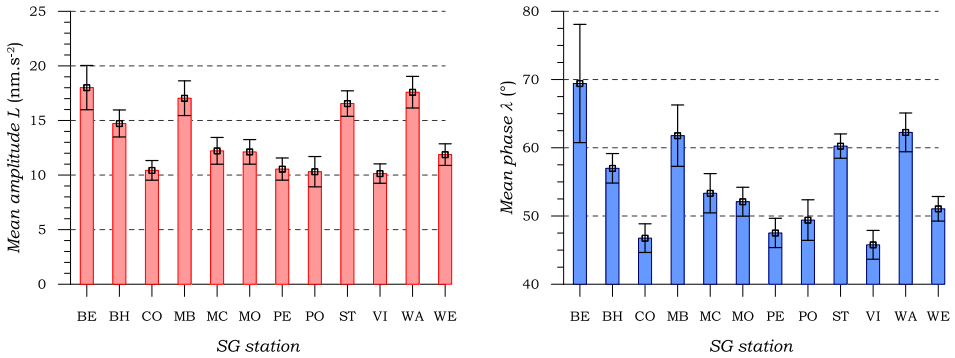


Fig. 6. Mean amplitudes (left) and phases (right) of M_2 ocean loading vector from the model predictions at different SG stations. Error bars show discrepancies between used ocean loading models (both multiplied by a factor of 5).

between 1.3 to $1.5 \text{ nm}\cdot\text{s}^{-2}$ and the phases between 140° and 150° . The load amplitudes of M_2 (10.0 to $18.0 \text{ nm}\cdot\text{s}^{-2}$) are about 10 times higher and phases vary between 46° and 70° . This is typical for the selected region as pointed out by *Baker and Bos (2003)* and confirms that the harmonic O_1 is especially well suited for validating body tide models. Note that standard deviations of mean amplitudes and phases in Fig. 6 have been multiplied by a factor of 5 for clarity reasons. Larger standard deviations are typical for SG stations near the coastal lines, e.g. Brussels, Membach, Medicina or Walferdange. It indicates the spatial resolution problem of global models in coastal areas with a strong ocean influence where such modeling is very complicated. In the case of intra-continental stations, e.g. Pecny, Vienna, Conrad and Wettzell, the discrepancies between the studied models are minimal.

5. Corrected tidal parameters

The ocean loading-corrected tidal parameters were calculated according to Eq. (1) taking eight ocean loading models into account as described in the previous section. Detailed formulas can be found in *Neumeyer (2010)* or *Sun et al. (1999)*. The numbers for O_1 and M_2 are presented in Tables 3

Table 3. Ocean loading-corrected amplitude factors δ_c and phases α_c derived from 8 ocean loading models for O₁. MEAN₁ values indicate the average of one model at 11 stations, while MEAN₂ represent the average of 8 models at one station (σ_1 is the standard deviation of MEAN₁ and σ_2 the standard deviation of MEAN₂)

SG	CSR4.0	GOT4.7	TPX07.2	FES2004	EOT08a	DTU10	HAMTIDE	EOT11a	MEAN ₂	STD ₂
	δ_c/α_c (-)/(°)									(-)/(°)
BE*	1.1595 <i>-0.01</i>	1.1564 <i>-0.09</i>	1.1563 <i>-0.07</i>	1.1567 <i>-0.07</i>	1.1565 <i>-0.11</i>	1.1562 <i>-0.10</i>	1.1567 <i>-0.09</i>	1.1564 <i>-0.10</i>	1.1568 -0.08	0.0011 <i>0.03</i>
BH	1.1539 <i>-0.03</i>	1.1540 <i>-0.03</i>	1.1538 <i>-0.02</i>	1.1542 <i>-0.01</i>	1.1540 <i>-0.04</i>	1.1537 <i>-0.04</i>	1.1542 <i>-0.02</i>	1.1539 <i>-0.04</i>	1.1540 -0.03	0.0002 <i>0.01</i>
CO	1.1531 <i>-0.02</i>	1.1534 <i>-0.02</i>	1.1534 <i>-0.01</i>	1.1536 <i>0.00</i>	1.1533 <i>-0.02</i>	1.1531 <i>-0.02</i>	1.1535 <i>-0.01</i>	1.1532 <i>-0.02</i>	1.1533 -0.01	0.0002 <i>0.01</i>
MB	1.1532 <i>-0.04</i>	1.1530 <i>-0.03</i>	1.1527 <i>-0.02</i>	1.1533 <i>-0.01</i>	1.1530 <i>-0.05</i>	1.1528 <i>-0.05</i>	1.1533 <i>-0.03</i>	1.1530 <i>-0.05</i>	1.1530 -0.03	0.0002 <i>0.01</i>
MC	1.1524 <i>0.08</i>	1.1524 <i>-0.01</i>	1.1522 <i>0.02</i>	1.1526 <i>-0.01</i>	1.1524 <i>-0.02</i>	1.1521 <i>-0.03</i>	1.1525 <i>-0.01</i>	1.1522 <i>-0.03</i>	1.1524 0.00	0.0002 <i>0.04</i>
MO	1.1526 <i>-0.01</i>	1.1527 <i>-0.01</i>	1.1527 <i>0.00</i>	1.1531 <i>0.00</i>	1.1527 <i>0.00</i>	1.1525 <i>-0.02</i>	1.1529 <i>-0.01</i>	1.1526 <i>-0.02</i>	1.1527 -0.01	0.0002 <i>0.01</i>
PE	1.1535 <i>-0.02</i>	1.1537 <i>-0.01</i>	1.1537 <i>-0.01</i>	1.1540 <i>0.00</i>	1.1536 <i>-0.02</i>	1.1535 <i>-0.02</i>	1.1539 <i>-0.01</i>	1.1535 <i>-0.02</i>	1.1537 -0.01	0.0002 <i>0.01</i>
PO	1.1536 <i>-0.01</i>	1.1536 <i>-0.04</i>	1.1535 <i>-0.02</i>	1.1541 <i>-0.01</i>	1.1536 <i>-0.04</i>	1.1534 <i>-0.04</i>	1.1539 <i>-0.03</i>	1.1535 <i>-0.04</i>	1.1536 -0.03	0.0002 <i>0.01</i>
ST	1.1531 <i>-0.03</i>	1.1533 <i>-0.03</i>	1.1532 <i>-0.02</i>	1.1535 <i>-0.01</i>	1.1532 <i>-0.03</i>	1.1530 <i>-0.04</i>	1.1534 <i>-0.02</i>	1.1532 <i>-0.04</i>	1.1532 -0.03	0.0002 <i>0.01</i>
VI	1.1531 <i>-0.01</i>	1.1534 <i>-0.01</i>	1.1534 <i>-0.01</i>	1.1536 <i>0.00</i>	1.1533 <i>-0.01</i>	1.1531 <i>-0.02</i>	1.1535 <i>-0.01</i>	1.1532 <i>-0.02</i>	1.1533 -0.01	0.0002 <i>0.01</i>
WA	1.1529 <i>-0.04</i>	1.1529 <i>-0.03</i>	1.1528 <i>-0.02</i>	1.1532 <i>-0.02</i>	1.1530 <i>-0.04</i>	1.1527 <i>-0.04</i>	1.1531 <i>-0.03</i>	1.1529 <i>-0.04</i>	1.1529 -0.03	0.0002 <i>0.01</i>
WE	1.1533 <i>-0.02</i>	1.1535 <i>-0.01</i>	1.1535 <i>-0.01</i>	1.1537 <i>0.00</i>	1.1534 <i>-0.02</i>	1.1532 <i>-0.02</i>	1.1536 <i>-0.01</i>	1.1533 <i>-0.02</i>	1.1534 -0.01	0.0002 <i>0.01</i>
MEAN ₁	1.1531 -0.01	1.1533 -0.02	1.1532 -0.01	1.1535 -0.01	1.1532 -0.03	1.1530 -0.03	1.1534 -0.02	1.1531 -0.03	1.1532 -0.02	σ_1 0.0002 <i>0.01</i>
STD ₁	0.0004 <i>0.03</i>	0.0004 <i>0.01</i>	0.0005 <i>0.01</i>	0.0005 <i>0.01</i>	0.0004 <i>0.01</i>	0.0005 <i>0.01</i>	0.0005 <i>0.01</i>	0.0005 <i>0.01</i>	σ_2 0.0005 0.01	

* not used in means

and 4. Two different comparisons of corrected amplitude factors δ_c and phase differences α_c were performed.

The first comparison is investigating the accuracy of ocean loading correction applied to the observed tidal parameters. For each station the mean

Table 4. Ocean loading-corrected amplitude factors δ_c and phases α_c derived from 8 ocean loading models for M_2 . $MEAN_1$ values indicate the average of one model at 11 stations, while $MEAN_2$ represent the average of 8 models at one station (σ_1 is the standard deviation of $MEAN_1$ and σ_2 the standard deviation of $MEAN_2$)

SG	CSR4.0	GOT4.7	TPXO7.2	FES2004	EOT08a	DTU10	HAMTIDE	EOT11a	MEAN ₂	STD ₂
	$\delta_c/\alpha_c (-)/(^{\circ})$									$(-)/(^{\circ})$
BE*	1.1596 <i>0.06</i>	1.1584 <i>-0.07</i>	1.1639 <i>0.04</i>	1.1624 <i>0.08</i>	1.1633 <i>-0.01</i>	1.1631 <i>-0.03</i>	1.1625 <i>-0.06</i>	1.1632 <i>-0.04</i>	1.1620 0.00	0.0020 <i>0.06</i>
BH	1.1620 <i>0.07</i>	1.1607 <i>0.00</i>	1.1620 <i>0.03</i>	1.1629 <i>0.09</i>	1.1623 <i>0.04</i>	1.1620 <i>0.03</i>	1.1620 <i>0.00</i>	1.1623 <i>0.02</i>	1.1620 0.04	0.0006 <i>0.03</i>
CO	1.1621 <i>0.03</i>	1.1613 <i>0.00</i>	1.1616 <i>0.00</i>	1.1630 <i>0.04</i>	1.1622 <i>0.02</i>	1.1618 <i>0.02</i>	1.1620 <i>0.00</i>	1.1620 <i>0.01</i>	1.1620 0.02	0.0005 <i>0.01</i>
MB	1.1594 <i>0.09</i>	1.1579 <i>-0.01</i>	1.1611 <i>0.03</i>	1.1608 <i>0.12</i>	1.1608 <i>0.05</i>	1.1605 <i>0.04</i>	1.1604 <i>0.00</i>	1.1608 <i>0.02</i>	1.1602 0.04	0.0010 <i>0.04</i>
MC	1.1611 <i>0.01</i>	1.1607 <i>-0.01</i>	1.1609 <i>-0.06</i>	1.1623 <i>0.03</i>	1.1615 <i>0.01</i>	1.1611 <i>0.02</i>	1.1613 <i>-0.01</i>	1.1613 <i>0.00</i>	1.1613 0.00	0.0005 <i>0.03</i>
MO	1.1606 <i>0.07</i>	1.1594 <i>0.02</i>	1.1604 <i>0.06</i>	1.1617 <i>0.10</i>	1.1607 <i>0.05</i>	1.1604 <i>0.05</i>	1.1604 <i>0.03</i>	1.1606 <i>0.03</i>	1.1605 0.05	0.0006 <i>0.03</i>
PE	1.1621 <i>0.04</i>	1.1610 <i>0.01</i>	1.1616 <i>0.02</i>	1.1631 <i>0.06</i>	1.1621 <i>0.03</i>	1.1617 <i>0.03</i>	1.1618 <i>0.01</i>	1.1619 <i>0.01</i>	1.1619 0.03	0.0006 <i>0.02</i>
PO	1.1622 <i>0.02</i>	1.1610 <i>-0.02</i>	1.1622 <i>0.02</i>	1.1641 <i>0.06</i>	1.1624 <i>0.00</i>	1.1620 <i>0.00</i>	1.1618 <i>-0.03</i>	1.1622 <i>-0.02</i>	1.1622 0.00	0.0009 <i>0.03</i>
ST	1.1616 <i>0.07</i>	1.1601 <i>0.00</i>	1.1613 <i>0.01</i>	1.1618 <i>0.08</i>	1.1615 <i>0.04</i>	1.1611 <i>0.04</i>	1.1614 <i>0.01</i>	1.1614 <i>0.02</i>	1.1613 0.04	0.0005 <i>0.03</i>
VI	1.1621 <i>0.03</i>	1.1612 <i>0.01</i>	1.1615 <i>0.01</i>	1.1629 <i>0.05</i>	1.1622 <i>0.03</i>	1.1618 <i>0.03</i>	1.1619 <i>0.01</i>	1.1620 <i>0.01</i>	1.1620 0.02	0.0005 <i>0.01</i>
WA	1.1604 <i>0.08</i>	1.1587 <i>-0.02</i>	1.1608 <i>0.00</i>	1.1608 <i>0.09</i>	1.1608 <i>0.04</i>	1.1605 <i>0.03</i>	1.1606 <i>0.00</i>	1.1608 <i>0.02</i>	1.1604 0.03	0.0007 <i>0.04</i>
WE	1.1621 <i>0.04</i>	1.1611 <i>0.00</i>	1.1617 <i>0.01</i>	1.1629 <i>0.06</i>	1.1621 <i>0.03</i>	1.1618 <i>0.03</i>	1.1619 <i>0.01</i>	1.1620 <i>0.01</i>	1.1619 0.03	0.0005 <i>0.02</i>
MEAN ₁	1.1614 0.05	1.1603 0.00	1.1614 0.01	1.1624 0.07	1.1617 0.03	1.1613 0.03	1.1614 0.00	1.1616 0.01	1.1614 0.03	σ_1 0.0006 0.03
STD ₁	0.0009 <i>0.03</i>	0.0011 <i>0.01</i>	0.0005 <i>0.03</i>	0.0010 <i>0.03</i>	0.0007 <i>0.02</i>	0.0006 <i>0.01</i>	0.0007 <i>0.01</i>	0.0006 <i>0.01</i>	σ_2 0.0007 0.02	

* not used in means

corrected amplitude factors and phase differences from all ocean loading models were computed. They are denoted as $MEAN_2$ in Tables 3 and 4. Their standard deviations STD_2 indicate the discrepancies between used models for individual SG sites. The largest deviations occur at coastal

stations like Brussels, Membach and Walferdange, essentially for M_2 due to higher amplitudes of the ocean loading vector. Therefore, the SG in Brussels was excluded from further analysis. To evaluate a global standard deviation σ_1 of the ocean loading correction for $n_1 = 8$ models the mean corrected tidal parameters $MEAN_1$ were calculated for each model considering all SG stations. The calibration errors affect the $MEAN_1$ values to the same extent in this case. This implies that a dispersion of $MEAN_1$ responds to the accuracy of ocean loading correction denoted as σ_1 . It is about 0.2 ‰ (O_1), 0.5 ‰ (M_2) for amplitude factors and 0.01° (O_1), 0.03° (M_2) for phase differences. The corrected phases should be negligible after removing the ocean loading effect. However, it is possible to observe their systematic behavior in both cases. For O_1 the phase differences are negative, while they are positive for M_2 .

Secondly, the influence of SG calibration errors on the tidal parameter estimation was studied. Discrepancies between $MEAN_2$ values are caused by calibration errors and different SG location. In this case, the ocean loading correction errors are less important and this allows for assessment of the global calibration error σ_2 affecting the tidal parameters for $n_2 = 11$ stations. It reaches 0.4 ‰ (O_1), 0.6 ‰ (M_2) for amplitude factors and 0.01° (O_1), 0.02° (M_2) for phase differences. Disparities of corrected amplitude factors due to different SG locations should not exceed 0.10 ‰ (0.0001) according to the latitude dependency of the theoretical models described in the next section. Finally, the standard deviations STD_1 express the discrepancies between SG stations for individual ocean loading models. Both global standard deviations may be regarded as the same, if their ratio σ_2/σ_1 is equal to $\sqrt{n_2/n_1} = 1.17$. This ratio is 2.65 (O_1) and 1.28 (M_2) for global standard deviations of amplitude factors. That means the SG calibration error causes higher uncertainty than the differences between the used ocean loading models. Regarding the phase differences, the ratios 1.11 (O_1) and 0.62 (M_2) are below the 1.17 criterion and suggests that instrumental phase lags are determined with higher accuracy than the ocean loading phases.

Figures 7 and 8 show the dependence of mean corrected tidal parameters on the SG location. The corrected amplitude factors of the intra-continental stations CO, PE, VI and WE differ by less than 0.4 ‰ for O_1 and less than 0.1 ‰ for M_2 . The closer the stations are located near the Atlantic, the smaller the factors. This is less pronounced for O_1 because the load vec-

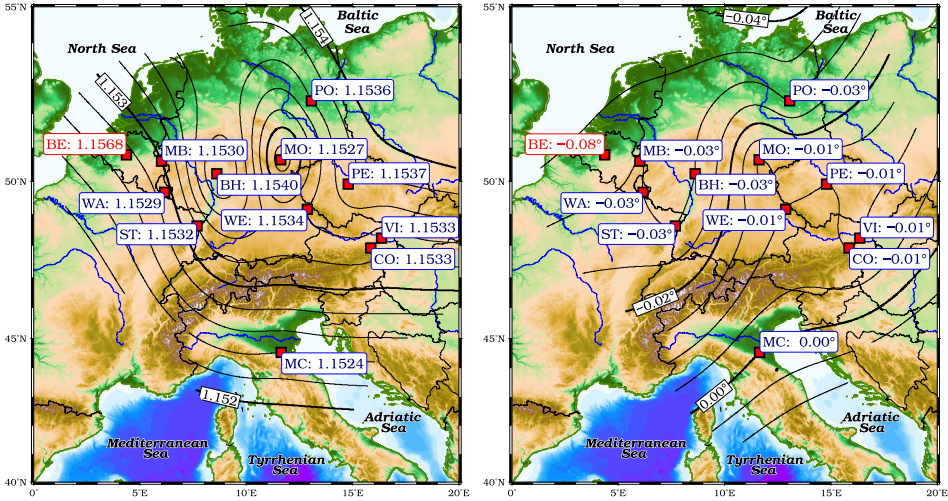


Fig. 7. Mean corrected amplitude factors (left) and phase differences (right) for O₁.

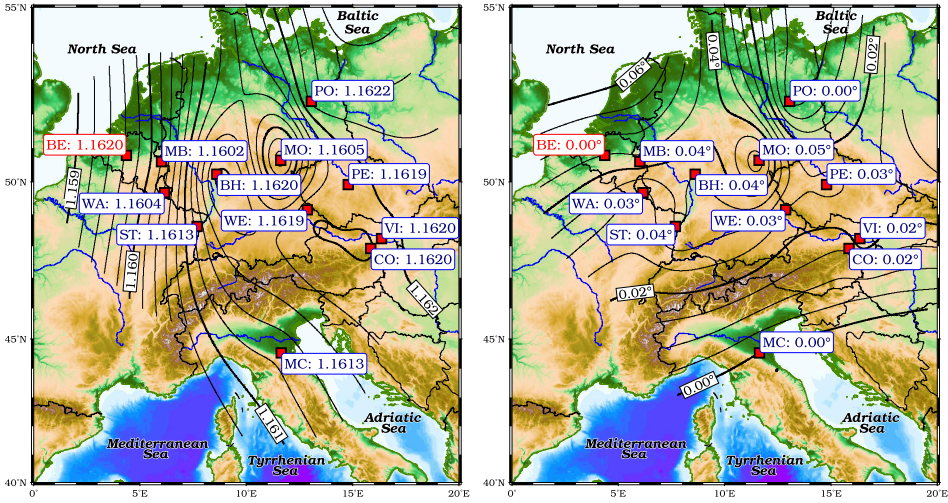


Fig. 8. Mean corrected amplitude factors (left) and phase differences (right) for M₂.

tor is smaller for O₁ than for M₂. Anomalous features at Moxa and Bad Homburg seem to appear similarly as in case of the uncorrected amplitude

factors (Figs. 3 and 4). Amplitude factors are too high at Bad Homburg and too low at Moxa both in the diurnal and semi-diurnal band while the M_2/O_1 ratio (see Table 5 and Fig. 9) does not show similarly pronounced anomalous features at these stations. Thus they are most likely due to calibration factor errors. The parameters in Brussels point to the resolution problem of global models close to the coastline and a poor calibration too.

As mentioned earlier, the corrected phases should be negligible after ocean loading correction. However, they are systematically changing from south-east to north-west (see Figs. 7 and 8). The phases are decreasing towards the North Sea for O_1 and increasing for M_2 , respectively. This systematic behavior can be explained by e.g. insufficient spatial resolution of global ocean loading models. A detailed error analysis has been done by *Bos and Baker (2005)*. There are exceptions in Brussels, Moxa and Potsdam. A curious situation is visible at the Medicina station close to the Mediterranean Sea. The corrected phase is 0.00° for both waves. However, higher standard deviations point to discrepancies between the used models. A similar case is M_2 at Brussels. In general, it should be considered that perhaps not all of the SGs are accurately calibrated with respect to their phase response. *Francis et al. (2011)* show that the time lags of three SGs differ by more than 1 s which corresponds to a phase difference of 0.004° in O_1 and 0.008° in M_2 . For sure, experimentally determined numbers are available for CO, MB, PE, ST, VI and WA.

At least in the investigated area, from these findings it appears that averaging various ocean loading models improves the reliability of ocean loading correction and of corrected tidal parameters, especially regarding the coastal SG stations. Good consistency between VI and CO confirms the stability of CT-025 calibration parameters after the transfer of the SG from Vienna to Conrad observatory (*Meurers, 2012*).

6. Comparison of results with body tide models

The corrected amplitude factors δ_c given in Tables 3 and 4 are now compared with the theoretical numbers δ_{th} provided by body tide models. We have used two recent body tide models for a rotating and non-spherical Earth (*Dehant et al., 1999*), hydrostatic/elastic DDW/He and non-hydrostatic/inelastic DDW/NHi models. The theoretical amplitude factors $\delta_{DDW/He}$ and

Table 5. Comparison of corrected (δ_c) and model amplitude factors ($\delta_{DDW/He}$, $\delta_{DDW/NHi}$). The model amplitude factors matching better the corrected ones are indicated in bold

SG	O_1					M_2					$\delta_c(M_2)/\delta_c(O_1)$ (-)
	δ_c (-)	DDW/He		DDW/NHi		δ_c (-)	DDW/He		DDW/NHi		
		δ_{th} (-)	$\Delta\delta$ (‰)	δ_{th} (-)	$\Delta\delta$ (‰)		δ_{th} (-)	$\Delta\delta$ (‰)	δ_{th} (-)	$\Delta\delta$ (‰)	
BE*	1.1568	1.15283	3.46	1.15431	2.18	1.1620	1.16049	1.32	1.16200	0.03	1.0045
BH	1.1540	1.15283	0.97	1.15431	-0.31	1.1620	1.16049	1.32	1.16199	0.03	1.0070
CO	1.1533	1.15282	0.43	1.15429	-0.85	1.1620	1.16047	1.32	1.16197	0.04	1.0075
MB	1.1530	1.15283	0.17	1.15431	-1.11	1.1602	1.16049	-0.25	1.16200	-1.54	1.0062
MC	1.1524	1.15281	-0.38	1.15427	-1.65	1.1613	1.16045	0.70	1.16193	-0.58	1.0077
MO	1.1527	1.15283	-0.10	1.15431	-1.38	1.1605	1.16049	0.03	1.16200	-1.27	1.0068
PE	1.1537	1.15283	0.72	1.15430	-0.56	1.1619	1.16049	1.21	1.16199	-0.08	1.0071
PO	1.1536	1.15284	0.70	1.15432	-0.59	1.1622	1.16051	1.49	1.16201	0.19	1.0075
ST	1.1532	1.15282	0.35	1.15430	-0.93	1.1613	1.16048	0.68	1.16197	-0.61	1.0070
VI	1.1533	1.15282	0.42	1.15429	-0.86	1.1620	1.16048	1.27	1.16197	-0.02	1.0075
WA	1.1529	1.15283	0.08	1.15430	-1.20	1.1604	1.16049	-0.06	1.16199	-1.35	1.0065
WE	1.1534	1.15282	0.53	1.15430	-0.74	1.1619	1.16048	1.25	1.16198	-0.04	1.0074
MEAN	1.1532	1.15283	0.35	1.15430	-0.93	1.1614	1.16048	0.81	1.16198	-0.48	1.0071

* not used in means

$\delta_{DDW/NHi}$ are listed in Table 5 and compared to the corrected factors δ_c . The difference between both models is about 1.3‰ which exceeds twice the global calibration σ_2 (0.6‰) of the M_2 corrected amplitude factors (see section 5). This allows establishing which body tide model fits better to the corrected amplitude factors. On the other hand, the latitude dependent theoretical amplitude factor variation between the northernmost SG in Potsdam and southernmost one in Medicina is less than 0.10‰.

We have calculated differences $\Delta\delta_{DDW/He}$ and $\Delta\delta_{DDW/NHi}$ as the corrected mean minus the theoretical amplitude factor for each SG station and body tide model. The deviations to the body tide models are of the same order of magnitude as the expected calibration accuracy (better than 1‰). However, a systematic behavior is clearly visible. For O_1 , the amplitude factors mostly deviate less from the DDW/He model than from the DDW/NHi model. The mean difference between $\bar{\delta}_c = 1.1532$ and $\bar{\delta}_{DDW/He} = 1.15283$ for all SG stations (except Brussels) is 0.35‰. Everywhere with except of Moxa (MO) and Medicina (MC) the corrected amplitude factors are slightly higher than predicted by DDW/He. Contrarily, for M_2 the DDW/NHi

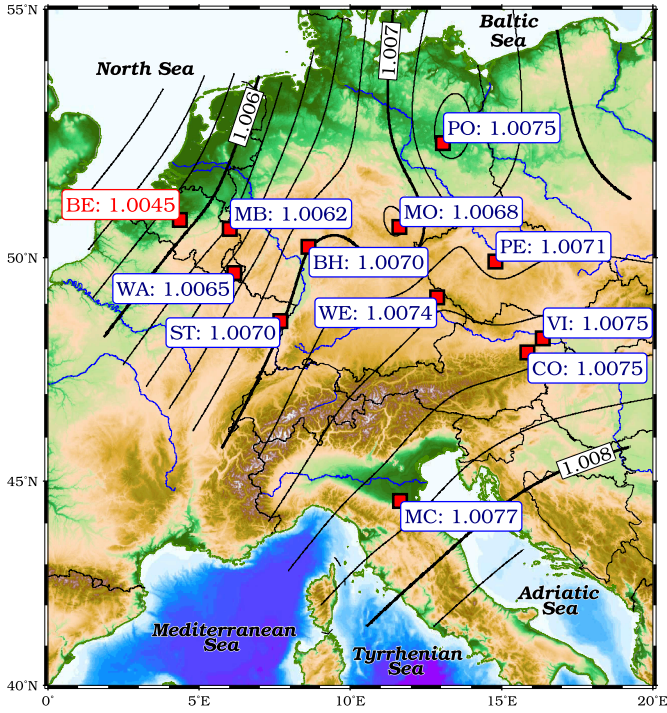


Fig. 9. The ratio $\delta_c(M_2)/\delta_c(O_1)$ of O_1 and M_2 corrected amplitude factors.

model matches the corrected amplitude factors better with a mean difference of -0.48‰ between $\bar{\delta}_c = 1.1614$ and $\bar{\delta}_{DDW/NHi} = 1.16198$. The corrected amplitude factors at the intra-continental stations CO, PE, VI and WE match the DDW/NHi model even better than 0.1‰ (average: -0.03‰). The deviations get larger and the M_2/O_1 ratio gets smaller when approaching the Atlantic sea (MB, ST, WA) due to the high load amplitude for M_2 and ocean model imperfections (see Fig. 9). This might also be the reason why MB and WA favor the DDW/He model (see Table 5). The numbers at BH and MO should be taken with care due to the suspected calibration problem. Generally, this result implies that the corrected amplitude factors are lying between the numbers derived from DDW/He and DDW/NHi models, but are close to the DDW/NHi model for M_2 .

7. Analysis of residual vectors

We can evaluate the efficiency of the used ocean loading models coupled with SGs’ calibration errors by analyzing the residual vectors \mathbf{B} and \mathbf{X} defined by Eqs. (2) and (3). The magnitude of the residual vector \mathbf{B} depends on the choice of the body tide model. Again, we have considered both DDW/He and DDW/NHi theoretical models. In Tables 6 and 7 the calculated amplitudes B and phases β are given. Observed amplitudes A are based on Tamura’s tidal potential (*Tamura, 1987*).

After subtraction of the ocean loading vector \mathbf{L} from the observed residual vector \mathbf{B} , we get the remaining residuals \mathbf{X} for each ocean loading model and station. Then, the mean amplitudes \bar{X} and phases $\bar{\chi}$ corresponding to single SG stations were calculated. Mean efficiency \bar{E} says to what extent the residual vector \mathbf{B} is reduced by the ocean loading correction considering a specific body tide model. The efficiency is about 80 % for O_1 and 97 % for M_2 . For O_1 , higher efficiency is achieved using the elastic DDW/He model

Table 6. Comparison of observed residuals \mathbf{B} and remaining residuals \mathbf{X} for O_1 together with the mean effectivity of ocean loading correction at SG stations

SG	A (nm.s ⁻²)	DDW/He					DDW/NHi				
		B	\bar{X}	β	$\bar{\chi}$	\bar{E}	B	\bar{X}	β	$\bar{\chi}$	\bar{E}
		(nm.s ⁻²)		(°)		(%)	(nm.s ⁻²)		(°)		(%)
BE*	351.08	0.41	1.33	65.18	-23.14	-229.4	0.46	0.95	127.33	-34.90	-104.9
BH	351.41	1.05	0.39	143.83	-26.10	63.2	1.44	0.20	154.51	-121.96	85.9
CO	355.36	1.18	0.19	144.78	-33.82	84.2	1.57	0.32	154.39	-163.71	79.8
MB	350.33	1.22	0.23	149.29	-74.21	81.4	1.63	0.45	157.42	-152.39	72.5
MC	356.71	1.58	0.23	146.69	-69.68	85.7	1.98	0.62	153.92	-86.93	68.5
MO	350.12	1.40	0.10	147.49	-98.61	92.6	1.80	0.49	155.23	-126.59	72.7
PE	352.11	1.13	0.27	139.38	-19.07	76.1	1.51	0.22	150.65	-115.01	85.6
PO	345.55	1.17	0.31	139.06	-35.78	73.6	1.53	0.27	150.07	-140.97	82.4
ST	354.14	1.27	0.22	157.03	-53.24	82.8	1.70	0.37	163.01	-153.23	78.1
VI	354.92	1.19	0.17	143.11	-25.07	86.0	1.58	0.31	153.05	-125.13	80.2
WA	352.21	1.37	0.21	157.33	-80.68	84.5	1.79	0.47	162.92	-154.86	73.6
WE	353.51	1.19	0.21	144.20	-25.77	81.9	1.58	0.28	153.89	-163.50	82.3
MEAN	352.40	1.25	0.23	146.56	-49.27	81.1	1.65	0.36	155.37	-136.75	78.3

* not used in means

Table 7. Comparison of observed residuals \mathbf{B} and remaining residuals \mathbf{X} for M_2 together with the mean effectivity of ocean loading correction at SG stations

SG	A (nm.s ⁻²)	DDW/He					DDW/NHi				
		B	\bar{X}	β	$\bar{\chi}$	\bar{E}	B	\bar{X}	β	$\bar{\chi}$	\bar{E}
		(nm.s ⁻²)	(nm.s ⁻²)	(°)		(%)	(nm.s ⁻²)	(nm.s ⁻²)	(°)		(%)
BE*	355.59	18.16	0.77	68.03	-4.38	95.7	17.99	0.58	69.36	-6.22	96.8
BH	365.61	15.17	0.53	55.95	19.44	96.5	14.91	0.28	57.42	44.18	98.1
CO	399.30	10.87	0.53	45.15	9.99	95.1	10.52	0.17	47.09	54.07	98.3
MB	359.53	17.23	0.39	62.43	54.00	97.7	17.02	0.65	63.79	108.03	96.2
MC	450.87	12.40	0.36	52.19	-5.06	97.1	12.07	0.34	54.32	-14.98	97.2
MO	358.31	12.38	0.36	52.97	94.09	97.1	12.11	0.59	54.68	143.77	95.2
PE	369.29	10.97	0.47	46.40	18.82	95.7	10.65	0.23	48.22	112.12	97.9
PO	332.28	10.65	0.51	47.50	-3.76	95.2	10.37	0.21	49.22	-13.72	98.0
ST	389.77	16.88	0.39	59.88	53.52	97.7	16.64	0.38	61.34	137.41	97.7
VI	394.29	10.59	0.52	44.44	16.32	95.1	10.24	0.21	46.39	102.56	98.0
WA	373.86	17.74	0.30	62.59	19.63	98.3	17.53	0.58	63.96	25.11	96.7
WE	381.24	12.30	0.50	49.85	17.34	95.9	11.99	0.22	51.61	106.99	98.2
MEAN	379.48	13.38	0.44	52.67	26.76	96.5	13.10	0.35	54.37	73.23	97.4

* not used in means

while for M_2 a better fit is obtained by applying the inelastic DDW/NHi model. It confirms the conclusions of the previous section. Larger remaining residuals \mathbf{X} are mainly caused by the inaccuracy of the global ocean loading models. However, residual phases $\bar{\chi}$ close to 0° or $\pm 180^\circ$ can also point to erroneous SG calibration factors (e.g. Brussels) or to body tide model misfit. On the other hand, phases close to $\pm 90^\circ$ can point to an instrumental phase lag or ocean loading phase problem. Actually, calibration and ocean loading model errors as well as the choice of a theoretical body tide model contribute to both components of the residual vectors. In Brussels, the ocean loading vector of O_1 overestimates the observed residuals \mathbf{B} several times, causing negative efficiency of -229% (DDW/He model).

In a similar way, the efficiency of single ocean loading models for all SGs was investigated. The results are shown in Fig. 10. The discrepancies between used models are negligible in most cases. It is obvious that an effective rate is depending on the body tide model as was already proved. However, the only cases where an ocean model reaches high efficiency us-

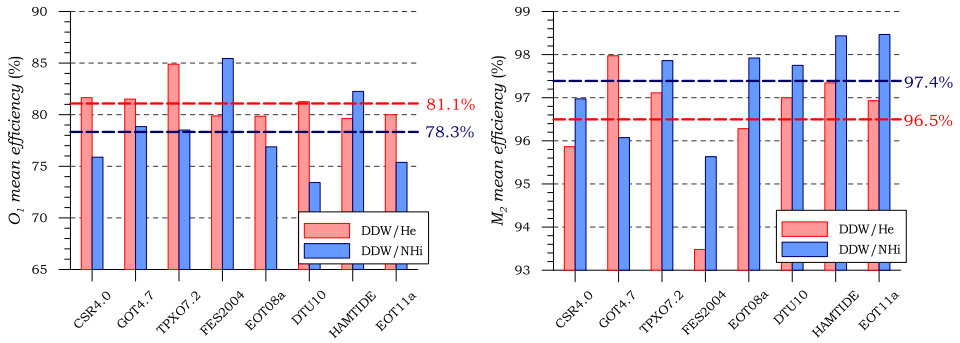


Fig. 10. Mean efficiency of ocean loading models for O_1 (left) and M_2 (right).

ing the same body tide model for both O_1 and M_2 is GOT4.7 (DDW/He model); FES2004 and HAMTIDE (DDW/NHi model).

Finally, it should be mentioned, that the analyzed time series have different lengths and refer to different epochs. Temporal variations of tidal parameter can occur e.g. due to non-linear ocean tides. In the study area, M_2 amplitude factor variations derived from 1 year intervals do not exceed $\pm 0.15\%$ (Meurers *et al.*, 2014). As we compare tidal parameters of much longer time series, the effect is certainly smaller than this.

8. Conclusions

A tidal analysis for 12 SGs located in the western and central parts of Europe was applied to 1 hour gravity data provided by the GGP database or SG owners. The influence of the decimation filters was studied using their filter transfer function. As the amplitude attenuation is -0.50% within the semi-diurnal tidal band for some filter methods, the observed amplitude factors were corrected according to amplitude response of the filter.

Corrected tidal parameters were calculated using eight recent ocean loading models (CSR4.0, GOT4.7, TPX07.2, FES2004, EOT08a, DTU10, HAMTIDE and EOT11a) calculated by the Free Ocean Tide Loading Provider (M.S. Bos and G.H. Scherneck). Models are still characterized by clear discrepancies between each other, especially for SGs near the coastal lines. In case of the intra-continental stations they do not differ significantly. The

comparison of corrected tidal parameters enables to derive a global accuracy of ocean loading correction that is about 0.2 ‰ (O_1), 0.5 ‰ (M_2) for amplitude factors and 0.01° (O_1), 0.03° (M_2) for phase differences. On the other hand, a global calibration error reaches 0.4 ‰ (O_1), 0.6 ‰ (M_2) and 0.01° (O_1), 0.02° (M_2), respectively. That means the SG calibration error causes higher uncertainty than the discrepancies between used ocean loading models. However, the instrumental phase lags are determined with higher accuracy than the ocean loading phases. The spatial distribution of corrected tidal parameters reveals both the calibration problem and the bias of phases after ocean loading correction.

Corrected amplitude factors were compared with theoretical values provided by the hydrostatic/elastic DDW/He and the non-hydrostatic/inelastic DDW/NHi body tide models (*Dehant et al., 1999*). Generally, the corrected amplitude factors of O_1 and M_2 are between the numbers of both body tide models. However, the O_1 amplitude factors fit closely to the DDW/He, while M_2 shows better fit to the DDW/NHi model with mean deviations of 0.35 ‰ and -0.48 ‰.

The analysis of residuals vectors allows assessing the efficiency of the used ocean loading models which is about 80 % (O_1) and 97 % (M_2). Higher efficiency is achieved for O_1 using the elastic DDW/He model and for M_2 using the inelastic DDW/NHi model.

Acknowledgments. The authors thank M. Van Camp (Royal Observatory of Belgium), B. Ducarme (Royal Observatory of Belgium), O. Francis (University of Luxembourg), M. Hendrickx (Royal Observatory of Belgium), T. Jahr (Institute of Geosciences, FSU Jena), J. Neumeier (GFZ Potsdam) and V. Pálinkáš (Research Institute of Geodesy, Topography and Cartography) for their interest and help with this study. Thanks also to H.G. Scherneck and M.S. Bos for the Free Ocean Tide Loading Provider. We gratefully acknowledge the reviewer comments and suggestions by M. Van Camp, which improved the paper essentially. This study is based on research carried out within the project SATIM-SG (Slovak University of Technology in Bratislava) and partially supported by the Slovak national project VEGA 1/1092/11. Cooperation with ZAMG (Central Institute of Meteorology and Geodynamics, Austria) is gratefully acknowledged.

References

- Baker T. F., Edge R. J., Jeffries G., 1991: Tidal gravity and ocean tide loading in Europe. *Geophysical Journal International*, **107**, 1, 1–11.

- Baker T. F., Curtis D. J., Dodson A. H., 1996: A new test of Earth tide models in central Europe. *Geophysical Research Letters*, **23**, 24, 3559–3562.
- Baker T. F., Bos M. S., 2003: Validating Earth and ocean tide models using tidal gravity measurements. *Geophysical Journal International*, **152**, 468–485.
- Bos M. S., Baker T. F., 2005: An estimate of the errors in gravity ocean tide loading computations. *Journal of Geodesy*, **79**, (1-3), 50–63.
- Cheng Y., Andersen O. B., 2010: Improvement in global ocean tide model in shallow water regions. OSTST Meeting, Oct. 18-22, Lisbon.
- Crossley D., Hinderer J., Casula G., Francis O., Hsu H. T., Imanishi Y., Jentzsch G., Kääriäinen J., Merriam J., Meurers B., Neumeyer J., Richter B., Shibuya K., Sato T., Van Dam T., 1999: Network of superconducting gravimeters benefits a number of disciplines. *Eos, Transactions, American Geophysical Union*, **80**, 11, 121–126. doi 10.1029/99E000079.
- Dehant V., Zschau J., 1989: The effect of mantle inelasticity on tidal gravity: a comparison between the spherical and the elliptical Earth model. *Geophysical Journal*, **97**, 549–555.
- Dehant V., Defraigne P., Wahr J., 1999: Tides for a convective Earth. *Journal of Geophysical Research B*, **104**, 1035–1058.
- Ducarme B., Somerhausen A., 1997: Tidal Gravity Recording at Brussels with a SCINTREX CG-3M Gravity Meter. *Marees Terrestres Bulletin d'Informations*, **126**, 9611–9634.
- Ducarme B., Sun H. P., Xu J. Q., 2002: New investigations of tidal gravity results from the GGP network. *Marees Terrestres Bulletin d'Informations*, **136**, 10761–10766.
- Ducarme B., Rosat S., Vandercoilden L., Xu J. Q., Sun H. P., 2009: European tidal gravity observations: comparison with Earth Tides models and estimation of the Free Core Nutation (FCN) parameters. *Observing our Changing Earth, International Association of Geodesy Symposia*, **133**, 3, 523–532. ISBN 978-3-540-85426-5.
- Eanes R. J., 1994: Diurnal and semidiurnal tides from TOPEX/POSEIDON altimetry. *Eos, Transactions American Geophysical Union*, **75**, 16, 108 p.
- Egbert G. D., Erofeeva L., 2002: Efficient inverse modeling of barotropic ocean tides. *Journal of Atmospheric and Oceanic Technology*, **19**, 2, 183–204.
- Francis O., 1997: Calibration of the C021 superconducting gravimeter in Membach (Belgium) using 47 days of absolute gravity measurements. In: *International association of Geodesy Symp., Gravity, Geoid and Marine Geodesy*, **117**, 212–219.
- Francis O., Melchior P., 1996: Tidal loading in south western Europe: A test area. *Geophysical Research Letters*, **23**, 2251–2254.
- Francis O., Lampitelli C., Klein G., Van Camp M., Pálinkáš V., 2011: Comparison between the Transfer Functions of three Superconducting Gravimeters. *Marees Terrestres Bulletin d'Informations*, **147**, 11857–11868.
- Lyard F., Lefèvre F., Letellier T., Francis O., 2006: Modelling the global ocean tides: a modern insight from FES2004. *Ocean Dynamics*, **56**, 394–415.
- Mathews P. M., 2001: Love numbers and gravimetric factors for diurnal tides. *Journal of the Geodetic Society of Japan*, **47**, 1, 231–236. ISSN 0038-0830.

- Melchior P., 1978: The tides of the planet earth. Pergamon Press, Oxford, 607 p.
- Melchior P., 1994: A new data bank for tidal gravity measurements (DB 92). *Physics of the Earth and Planetary Interiors*, **82**, 125–155.
- Melchior P., Ducarme B., Francis O., 1996: The response of the Earth to tidal body forces described by second- and third-degree spherical harmonics as derived from a 12 year series of measurements with the superconducting gravimeter GWR/T3 in Brussels. *Physics of the Earth and Planetary Interiors*, **93**, 223–238.
- Meurers B., 2012: Superconducting Gravimeter Calibration by Colocated Gravity Observations: Results from GWR C025. *International Journal of Geophysics*, **2012**, 12 p. doi 10.1155/2012/954271.
- Meurers B., Van Camp M., Pálinkáš V., 2014: Temporal Variation of Tidal Parameters. EGU General Assembly 2014, *Geophysical Research Abstracts*, Vol. **16**, EGU2014-13225, 2014.
- Neumeier J., Pino J., Dierks O., Sun H. P., Pflug H., 2005: Improvement of ocean loading correction on gravity data with additional tide gauge measurements. *Journal of Geodynamics*, **40**, 104–111.
- Neumeier J., 2010: Superconducting Gravimetry. In: Xu G., *Sciences of Geodesy I, Advances and Future Directions*. Berlin: Springer-Verlag, 339–413. ISBN 978-3-642-11740-4.
- Ray R. D., 1999: A global ocean tide Model from TOPEX/POSEIDON Altimetry: GOT99.2. NASA Technical Memorandum 209478.
- Savcenko R., Bosch W., 2008: EOT08a, empirical ocean tide model from multi-mission satellite altimetry. Report No. 81. Deutsches Geodätisches Forschungsinstitut (DGFI), München.
- Savcenko R., Bosch W., 2011: EOT11a – a new tide model from Multi-Mission Altimetry. OSTST Meeting, Oct. 19-21, San Diego.
- Sun H. P., Hsu H., Luo S., Xu J., 1999: Study of the ocean tidal models using tidal gravity observations obtained with superconducting gravimeter. *Acta Geodetica et Cartographica Sinica*, **28**, 2, 64–71.
- Taguchi E., Stammer D., Zahel W., 2010: Estimation of deep ocean tidal energy dissipation based on the high-resolution data-assimilative HAMTIDE model. (to be submitted to *J. Geophys. Res.*).
- Tamura Y. A., 1987: Harmonic development of the tide-generating potential. *Marees Terrestres Bulletin d'Informations*, **99**, 6813–6855.
- Van Camp M., Vauterin P., Wenzel H. G., Schott P., Francis O., 2000: Accurate transfer function determination for superconducting gravimeters. *Geophysical Research Letters*, **27**, 37–40.
- Van Camp M., Vauterin P., 2005: Tsoft: graphical and interactive software for the analysis of time series and Earth tides. *Computers & Geosciences*, **31**, 5, 631–640. ISSN 00983004.
- Wenzel H. G., 1996: The nanogal software: Earth tide data processing package ETERNA 3.30. *Marees Terrestres Bulletin d'Informations*, **124**, 2, 9425–9439. ISSN 1340-4202.